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### **Ocean Optics Protocols For Satellite Ocean Color Sensor Validation, Revision 3, Volume 2**

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## Chapter 12

### Above-Water Radiance and Remote Sensing Reflectance Measurement and Analysis Protocols

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#### 12.1 INTRODUCTION

As an alternative to the in-water methods of Chapters 10 and 11, water-leaving radiance can be measured from the deck of a ship. A shipboard radiometer is used to measure radiance  $L_{\text{sfc}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o)$  emanating from the sea surface at zenith angle  $\theta$  (usually chosen between  $30^\circ$  and  $50^\circ$ ) and azimuth angle  $\phi$  (usually chosen between  $90^\circ$  and  $180^\circ$  away the sun's azimuth  $\phi_o$ ). In the convention used here, azimuth angles  $\phi$  are measured relative to the sun's azimuth, *i.e.*  $\phi_o = 0$ .

The surface radiance measured with a radiometer having a solid-angle field of view (FOV) of  $\Omega_{\text{FOV}}$  sr may be expressed, following Mobley (1999), as

$$L_{\text{sfc}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o) = L_{\text{w}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o) + \rho L_{\text{sky}}(\lambda, \theta_{\text{sky}}, \phi_{\text{sky}} \in \Omega'_{\text{FOV}}; \theta_o). \quad (12.1)$$

$L_{\text{w}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o)$  is water-leaving radiance centered at angles  $(\theta, \phi)$  and averaged over  $\Omega_{\text{FOV}}$  [as weighted by the radiometer's directional response function (see Chapter 5)].  $L_{\text{sky}}(\lambda, \theta_{\text{sky}}, \phi_{\text{sky}} \in \Omega'_{\text{FOV}}; \theta_o)$  is sky radiance measured with the radiometer looking upward at angles  $(\theta_{\text{sky}}, \phi_{\text{sky}})$ . In practice,  $\theta$  and  $\theta_{\text{sky}}$  are numerically equal angles in the nadir and zenith directions, respectively, and the sea and sky viewing azimuths  $\phi = \phi_{\text{sky}}$ . The reflectance factor  $\rho$  is operationally defined as the total skylight actually reflected from the wave-roughened sea surface into direction  $(\theta, \phi)$  divided by sky radiance measured with the radiometer from direction  $(\theta_{\text{sky}}, \phi_{\text{sky}})$ , both quantities being averaged over  $\Omega_{\text{FOV}}$  (Mobley 1999). Remote sensing reflectance is then determined, using water-leaving radiance calculated from (12.1), as

$$R_{\text{RS}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o) = \frac{L_{\text{w}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o)}{E_{\text{s}}(\lambda; \theta_o)}, \quad (12.2)$$

where  $E_{\text{s}}(\lambda; \theta_o)$  is incident spectral irradiance measured above the sea surface. All of the above variables vary with solar zenith angle  $\theta_o$ .

A simplified notation is used in Chapters 10 and 11 (and elsewhere in the protocols) when discussing water leaving radiance  $L_{\text{w}}(\lambda)$  and remote sensing reflectance  $R_{\text{RS}}(\lambda)$  derived from in-water profile measurements of  $L_{\text{u}}(z, \lambda)$ . Because  $L_{\text{u}}(z, \lambda)$  is measured viewing the nadir direction,  $L_{\text{w}}(\lambda)$  represents radiance leaving the surface in the zenith direction  $(\theta, \phi) = (0^\circ, 0^\circ)$ . Therefore,  $L_{\text{w}}(\lambda)$  in Chapter 11 corresponds to  $L(\lambda, 0, 0 \in \Omega; \theta_o)$ , and  $R_{\text{RS}}(\lambda)$  to  $R(\lambda, 0, 0 \in \Omega; \theta_o)$ , in the present notation.